

A satellite image of Earth showing a vast expanse of white and grey cloud cover over a dark blue ocean. The horizon of the Earth is visible at the top of the frame, with a thin blue line representing the atmosphere against the black background of space.

# Quantifying radiative perturbations from observations and application to the short-term water vapor feedback

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- Mostly motivated by the need evaluate feedbacks in GCMs
  - ➊ Partial radiative perturbation (PRP) calculations: calculations that substitutes one-at-a-time a variable from the perturb climate into the control climate.  
(Wetherald and Manabe 1988)
  - ➋ Radiative kernels: separates the radiative response and the perturbations  
(Held and Soden 2000; Soden and Held 2006; Soden et al. 2008)
    - Fewer computations: a single radiative calculation can be consistently applied across different climate models.
    - Radiative kernel have become an indispensable tool for GCM feedback studies

# Observed radiative perturbations

- GCM-derived radiative kernels also been applied to observations in attempt to constrain climate feedbacks (Dessler 2008, 2010, 2013; Dessler and Wong 2009; Masters 2012; Dessler and Loeb 2013; Gordon et al. 2013, Zhou et al. 2013, 2014; Ceppi 2016)
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- But, there is a concerted effort to determine how observations can best be used to understand climate feedbacks and sensitivity (Loeb et al. 2016; Forster 2016)
  
- **Limited comparisons of GCM vs. observational-based kernels**  
(different GCM kernels show relatively small differences)
- **Would be attractive to perform calculations based purely on observations**

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- Examine feedbacks in the observational record within a consistent framework and construct observationally-based radiative kernels
- Provide general insight into the variability of the radiation budget
  - Useful where radiative effects are correlated among parameters.
    - E.g. clouds: compliment to the cloud radiative effect (CRE, total - clear-sky fluxes)

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- Easy to incorporate multiple datasets for each input

# PRP calculations

## PRP calculations

$$\partial F_{\Delta x}^f = F(x, y_1, \dots, y_N) - F(\bar{x}, y_1, \dots, y_N) \quad (1)$$

- Flux ( $F$ ) difference of monthly means ( $x, y$ ) and climatological monthly means ( $\bar{x}, \bar{y}$ )



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Can also compute the same thing relative to a different base state:

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- From monthly-mean inputs, climatologies are constructed and the variables combined to make the 4 sets of inputs → Fu-Liou radiative model

# Inputs (mostly) from CERES datasets

- Clear-sky: **GEOS** | **AIRS** (AIRX3STM)  
*temperature, water vapor, ozone, skin temperature*
- Clouds: **SYN** | **C3M**  
*fraction, base, top, phase, optical depth, size*
- Aerosol: **MATCH**  
*optical depth, vertical distribution, type*
- Gases: **AIRS** (AIRS3C2M/AIRX3STM)  
*carbon dioxide, methane*
- Gases: **NOAA ESRL** (global means)  
*nitrous oxide, CFC-11, CFC-12, HCFC-22*
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*parameterization spectral dependence*
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**Focus here on 13 years of  
SYN + GEOS/AIRS  
(Sept. 2002 – Aug. 2015)**



# Validation

❶ Perturbations add linearly:  $\Delta F = F(y_1, \dots, y_N) - F(\overline{y}_1, \dots, \overline{y}_N) \approx \sum_{i=1}^n \partial F_{\Delta x_i}$

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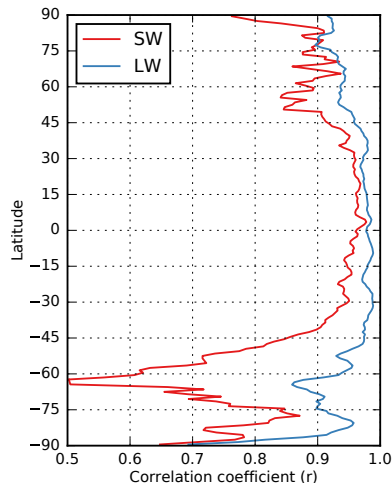
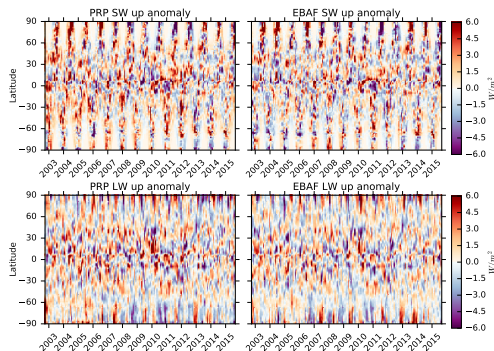
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- ❸ Reproduce the variability as observed by CERES
  - Time series well correlated over much of the globe (expect  $\sim 60^\circ\text{S}$ )



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	Feedback	Uncertainty	Dataset
Forster and Collins (2004)	1.6	0.9–2.5	NVAP, MLS
Dessler et al. (2008)	2.04	0.94–2.69	AIRS
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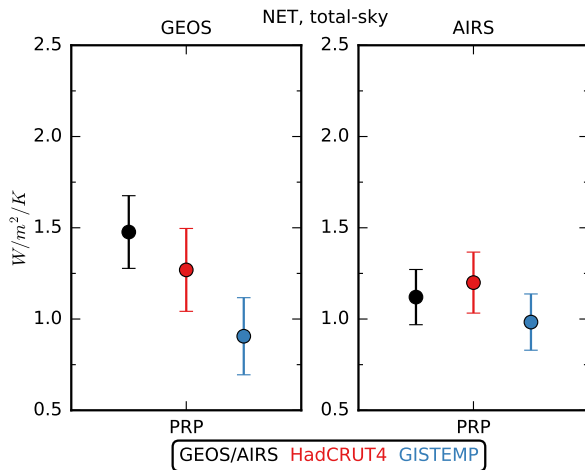
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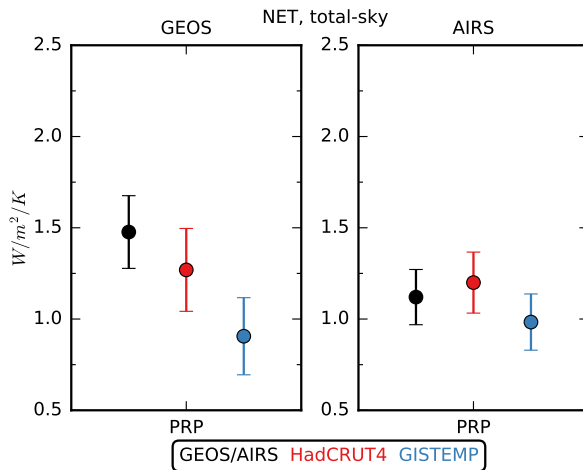
**Revisit this estimate with our PRP calculations, longer datasets (13 yrs), and more surface temperature datasets**

- Compute monthly perturbation to TOA flux caused by tropospheric water vapor:  $\Delta F$
- Feedback = slope of least-squares fit of  $\Delta F$  and surface temperature anomaly

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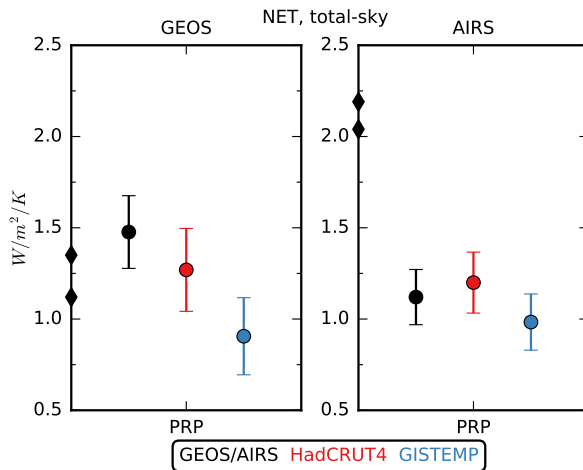


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- AIRS feedback is generally smaller
- AIRS is “tighter”: less sensitivity to  $T_{sfc}$  dataset and better fits

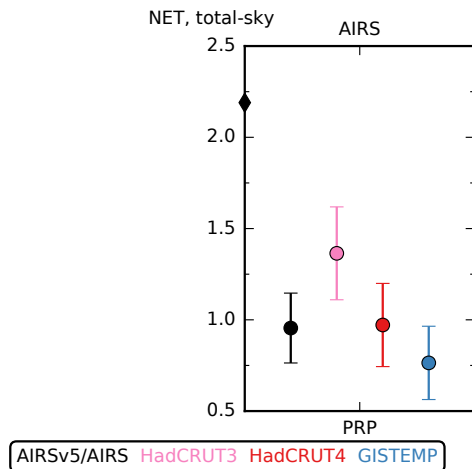
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- GEOS in agreement with previous reanalysis estimates
- **AIRS feedback nearly half that of previous estimates**
  - D08/G13: different period of data, older versions of datasets, kernels for getting  $\Delta F$

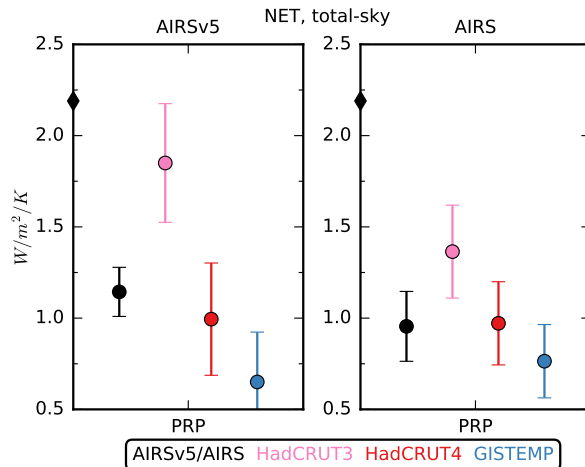
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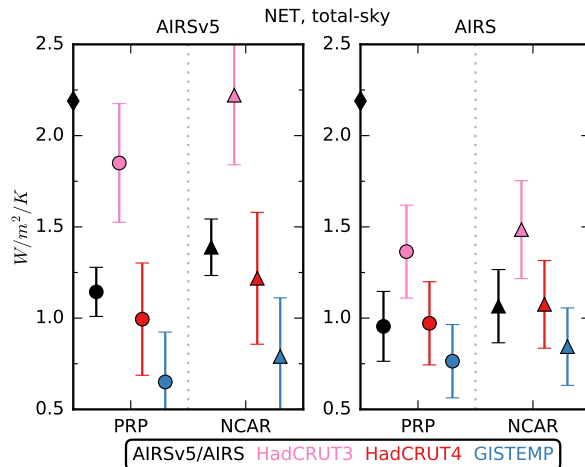
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- AIRSv5 + HadCRUT3 = largest feedback
- AIRSv5 + HadCRUT3 + NCAR kernel (Shell et al. 2008)  $\approx$  G13
  - AIRSv6/HadCRUT4 increase yield/coverage

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- Older water vapor feedback estimates using AIRSv5 + HadCRUT3 nearly 2x larger: mostly due to dataset updates (also differences from kernel, length of data)



## Looking forward:

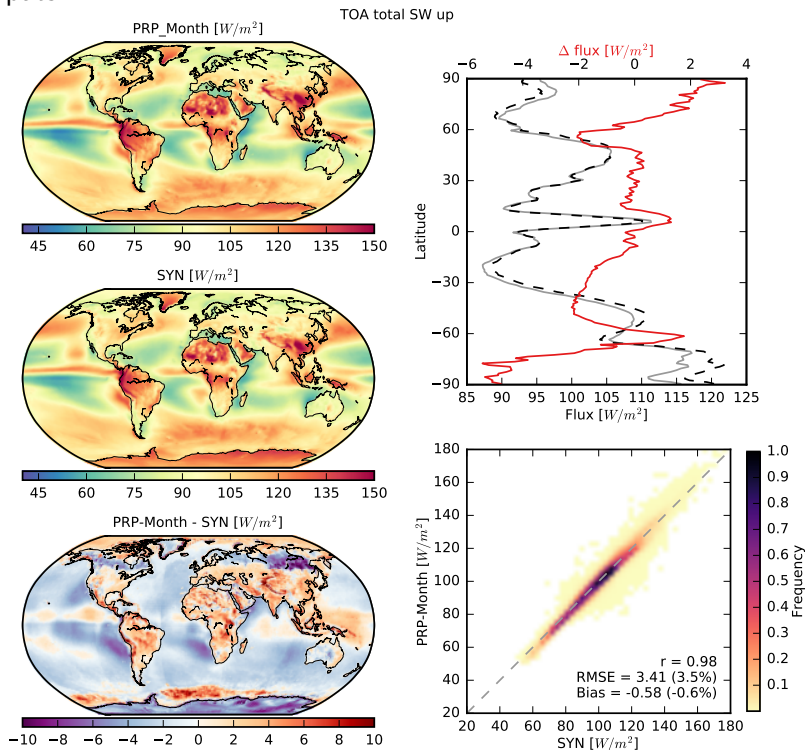
- AIRS v6.1: overhaul of water vapor retrievals
- Interesting to perform calculations with active sensor clouds (C3M)  
(likely to decrease water vapor feedback a bit)
- CERES-optimized radiative kernels: “two-pass” calculation to require that the sum of individual flux anomalies match CERES  
(similar to Sanderson and Shell 2012)



## Use of monthly mean inputs

Fluxes from  
monthly-mean inputs  
(**PRP\_Month**) vs.  
average fluxes computed  
in **SYN**

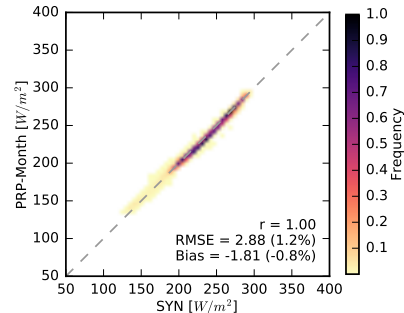
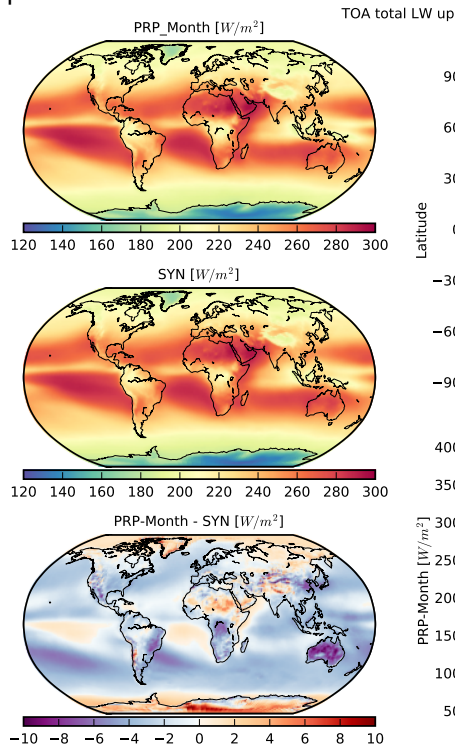
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(SW, LW)
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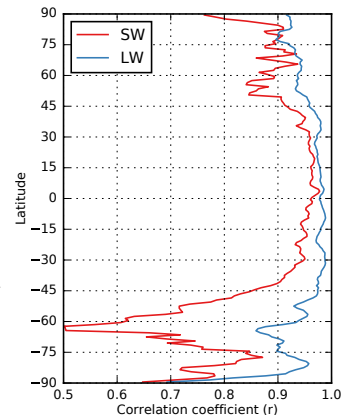
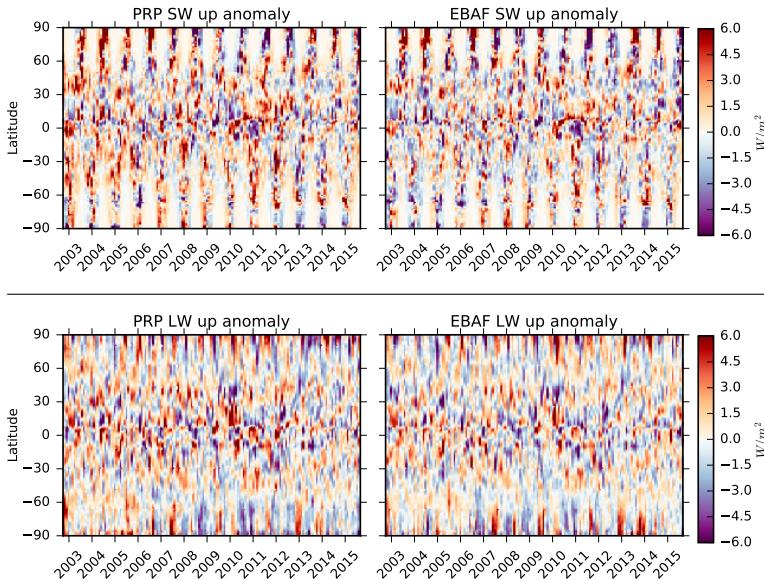
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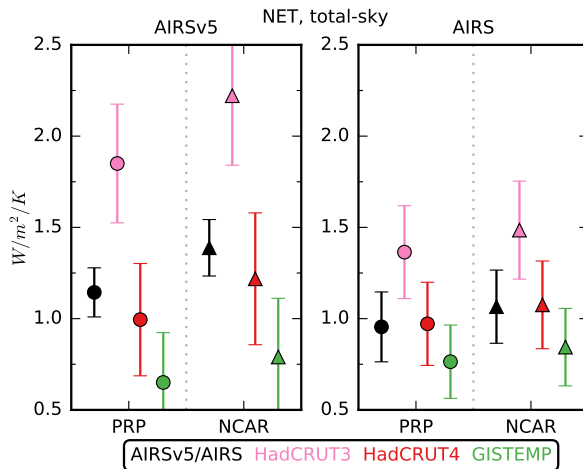


# Comparison to CERES (EBAF)



- Time series well correlated over much of the globe
- 60°S: ??

# Comparison to NCAR kernel

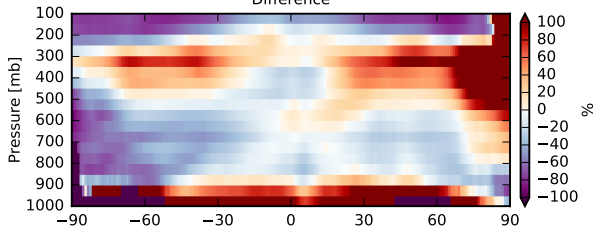
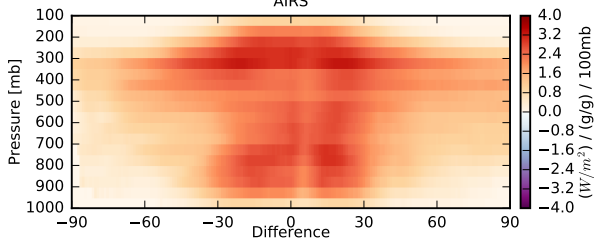
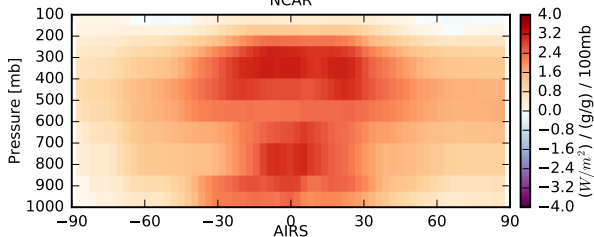


- Using NCAR kernel give a (slightly) larger feedback
- Kernel differences: (not shown)
  - Zonal / vertical means:  $\sim 10\%$   
(about the same as among different GCMs (Soden et al. 2008))
  - Local differences:  $\sim 20\text{--}40\%$   
(larger than among different GCMs (Soden et al. 2008))

# Comparison to NCAR kernel (Shell et al. 2008)

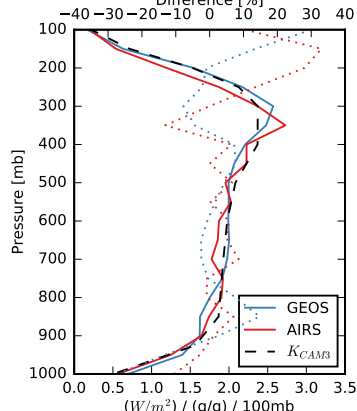
NET total-sky kernel

NCAR

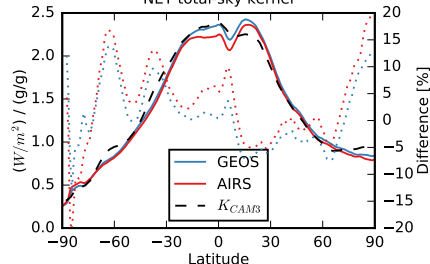


NET total-sky kernel

Difference [%]



NET total-sky kernel



# Short vs. Long-term water vapor feedback

- Gordon et al. (2013): short-term feedbacks in CMIP3 models converge to 15% of their long-term value after 25 years